



Updated aerosol  
module and its  
application for SOA

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# Updated aerosol module and its application to simulate secondary organic aerosols during IMPACT campaign May 2008

Y. P. Li<sup>1,2,3</sup>, H. Elbern<sup>2,3</sup>, K. D. Lu<sup>1</sup>, E. Friese<sup>2</sup>, A. Kiendler-Scharr<sup>3</sup>, Th. F. Mentel<sup>3</sup>, X. S. Wang<sup>1</sup>, A. Wahner<sup>3</sup>, and Y. H. Zhang<sup>1</sup>

<sup>1</sup>State Key Joint Laboratory of Environmental Simulation and Pollution Control, College of Environmental Sciences and Engineering, Peking University, Beijing, China

<sup>2</sup>Rhenish Institute for Environmental Research at the University of Cologne, Cologne, Germany

<sup>3</sup>Institut für Energie- und Klimaforschung Troposphäre (IEK-8), Forschungszentrum Jülich, Jülich, Germany

Received: 20 February 2013 – Accepted: 21 February 2013 – Published: 6 March 2013

Correspondence to: Y. H. Zhang (yhzhang@pku.edu.cn)

Published by Copernicus Publications on behalf of the European Geosciences Union.

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## Abstract

The formation of Secondary organic aerosol (SOA) was simulated with the Secondary ORGanic Aerosol Model (SORGAM) by a classical gas-particle partitioning concept, using the two-product model approach, which is widely used in chemical transport models. In this study, we extensively updated SORGAM including three major modifications: firstly, we derived temperature dependence functions of the SOA yields for aromatics and biogenic VOCs, based on recent chamber studies within a sophisticated mathematic optimization framework; secondly, we implemented the SOA formation pathways from photo oxidation (OH initiated) of isoprene; thirdly, we implemented the SOA formation channel from NO<sub>3</sub>-initiated oxidation of reactive biogenic hydrocarbons (isoprene and monoterpenes). The temperature dependence functions of the SOA yields were validated against available chamber experiments. Moreover, the whole updated SORGAM module was validated against ambient SOA observations represented by the summed oxygenated organic aerosol (OOA) concentrations abstracted from Aerosol Mass Spectrometer (AMS) measurements at a rural site near Rotterdam, the Netherlands, performed during the IMPACT campaign in May 2008. In this case, we embedded both the original and the updated SORGAM module into the EUROpean Air pollution and Dispersion-Inverse Model (EURAD-IM), which showed general good agreements with the observed meteorological parameters and several secondary products such as O<sub>3</sub>, sulfate and nitrate. With the updated SORGAM module, the EURAD-IM model also captured the observed SOA concentrations reasonably well especially those during nighttime. In contrast, the EURAD-IM model before update underestimated the observations by a factor of up to 5. The large improvements of the modeled SOA concentrations by updated SORGAM were attributed to the mentioned three modifications. Embedding the temperature dependence functions of the SOA yields, including the new pathways from isoprene photo oxidations, and switching on the SOA formation from NO<sub>3</sub> initiated biogenic VOCs oxidations contributed to this enhancement by 10 %, 22 % and 47 %, respectively. However, the EURAD-IM model

ACPD

13, 5961–6005, 2013

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with updated SORGAM still clearly underestimated the afternoon SOA observations up to a factor of two. More work such as to improve the simulated OH concentrations under high VOCs and low NO<sub>x</sub> concentrations, further including the SOA formation from semi-volatile organic compounds, the correct aging process of aerosols, oligomerization process and the influence on the biogenic SOA by the anthropogenic SOA, are still required to fill the gap.

## 1 Introduction

At local or regional scales, fine particle pollution is one of the central concerns of air pollution control due to its negative health impacts, degradation of visibility, and potentially enables active heterogeneous reactions. According to the available ambient measurements performed at urban, urban downwind area and rural locations, organic aerosol comprises a major fraction of submicron particulates (Zhang et al., 2007); of which, secondary organic aerosol (SOA) is a significant portion (Hallquist et al., 2009; Heald et al., 2005). However, it is still a challenging work to simulate SOA using chemical transport models, as both the diagnosis and prediction of the fine particle pollutions from local to global scales are still uncertain, yet form a basis for projecting climate change and regional air pollution control.

Several chemical transport models (CTM), such as GEOS-CHEM, Community Multi-scale Air Quality (CMAQ), PM Comprehensive Air Quality Model (PMCAMx) (Fu et al., 2008; Carlton et al., 2010; Lane et al., 2008), were employed to simulate SOA formation. However, there was large uncertainty of current models to predict SOA in the atmosphere due to the large varieties of SOA precursors and processes involved in formation of SOA (Hallquist et al., 2009). A new study even indicated that the equilibrium partitioning method used in traditional models might be questionable at real atmospheric conditions (Perraud, et al., 2012). The widely used frameworks to simulate SOA formation processes include the two-product model approach (Odum et al., 1996), the methods based on volatility basis set (VBS, Donahue et al., 2006) such

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as the 2-D-VBS (Jimenez et al., 2009) and statistical oxidation model (SOM, Cappa et al., 2013), etc. Additionally in recent studies, other new frameworks such as the condensed aromatic kinetics (Kamens et al., 2011; Zhou et al., 2011) and the functional group oxidation model (Zhang et al., 2012) were developed as well.

The EUROpean Air pollution and Dispersion-Inverse Model (EURAD-IM, Elbern et al., 2007) is an Eulerian model running from local to continental scale. EURAD-IM is primarily used for the chemical weather forecast and advanced data assimilation studies over Europe. Previous studies on a high ozone episode (Monteiro et al., 2012) and a dust storm (Chervenkov and Jakobs, 2011) indicated good performances of EURAD-IM. Colette et al. (2011) compared six CTMs including BOLCHEM, CHIMERE, EMEP, EURAD, OSLOCTM2 and MOZART used for simulations of the European air quality. It was found that EURAD-IM was better in simulating NO<sub>2</sub> and PM<sub>10</sub> among all the six CTMs. The SOA module in EURAD, namely, Secondary ORGanic Aerosol Model (SORGAM, Schell et al., 2001), is based on the two-product approach, but the performance of the SOA simulations by EURAD-IM is rarely evaluated against European ambient measurements.

The knowledge of SOA formation was largely developed in the last two decades. New precursors or new pathways of the oxidation of known precursors were recognized and sequentially implemented into the model. The photo oxidation products of monoterpenes and aromatics were firstly considered and determined to contribute to the aerosol budget due to their relative large carbon skeleton (Kavouras et al., 1998; Odum et al., 1997). In contrast, isoprene, the most prominent ambient VOC (Kesselmeier and Staudt, 1999; Guenther et al., 1996; Guenther et al., 2006) was initially not considered to be a possible SOA precursor, since it was a small molecule (Pandis et al., 1991). However, interesting indications that isoprene could be an important SOA precursor started to appear when Claeys et al. (2004) detected two polyols containing isoprene skeleton from natural aerosols. Furthermore, followed by the parameterization given by Kroll et al. (2006), a global model study estimated that the simulated aerosol burden increased by a factor of two by considering the isoprene

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generated SOA formation (Henze and Seinfeld, 2006). In a recent kinetic study, Paulot et al. (2009) proposed that the produced isoprene epoxides and the subsequent formation of polymers on aerosols could be one important physical mechanism for the isoprene SOA formation. Moreover, recent studies also highlighted the importance of reactive nitrogen compounds in global SOA formation. For example, degradation products from the NO<sub>3</sub> radical initiated oxidation of VOCs was found to be very important for SOA formation during nighttime conditions (Fry et al., 2009, 2011; Rollins et al., 2009; Pye et al., 2010).

In addition to the chemistry, ambient conditions such as temperature and humidity can also strongly influence the SOA production rates as observed in a series of chamber experiments. Temperature was reported as an important factor (Chung and Seinfeld, 2002), influencing the SOA yield from a given gas phase precursor-oxidant combination. The influence of temperature was manifold as it can affect the gas-phase oxidation mechanism, thus shifting the product distribution, as well as the partitioning of vapors between the gas phase and the particulate phase. Photo-oxidation experiments in a temperature-controlled chamber showed that, for three aromatic hydrocarbons (toluene, m-xylene and 1,2,4-trimethylbenzene),  $\alpha$ -pinene and n-undecane (Takekawa et al., 2003), SOA yield at 283 K was approximately twice as much than at 303 K. On the other hand, at very low temperature (between 243 K and 273 K), positive temperature dependence was observed and ascribed to changes in chemical kinetics (Jonsson et al., 2008). For SOA formed from m-xylene, a transition point near 290–292 K was observed for SOA density, volatility, hygroscopicity and elemental composition, when the experimental temperatures were increased. This indicated that simply employing the enthalpy of vaporization was not sufficient in predicting temperature dependent SOA formation (Qi et al., 2010). In addition, the influence of humidity on SOA formation was shown in former works (Seinfeld et al., 2001; Seinfeld and Pankow, 2003). More SOA mass was produced from the same amount of  $\alpha$ -pinene in the humid cases than that under dry conditions, with the total yield increasing about four times at both high and

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low temperatures (Tillmann et al., 2010). Here reaction channels involving water led to formation of products with lower vapor pressures.

To transfer the chamber studies into model frameworks, parameterizations of SOA formation were needed. Chamber experiments on the effect of temperature on the SOA formation from ozonolysis of  $\alpha$ -pinene and limonene were reported by Saathoff et al. (2009). They used the COSIMA-SOA model (Naumann et al., 2003) to reproduce the observed SOA mass, number and size distributions by parameterizing the temperature dependence of mass stoichiometric and partitioning coefficients. Stanier et al. (2008) used experimental smog chamber data of  $\alpha$ -pinene ozonolysis to parameterize the aerosol mass fractions. Based on the theory of absorptive partitioning and the method of the volatility basis set (VBS), a nonlinear least squares fitting tool was employed to calculate the mass stoichiometric coefficient and the enthalpy of vaporization. However, the parameterizations in recent works (Saathoff et al., 2009, etc.) were usually derived from one group of experiments, which might not be representative. In some cases (Stanier et al., 2008, etc.) temperature sensitivity was considered based on multi-researches, but it did not give an explicit function to describe the relations among the researches. In most chemical transport models (Henze and Seinfeld, 2006; Carlton et al., 2010, etc.), mass stoichiometric and partitioning coefficients were set as constants at a reference temperature (298 K), and the enthalpy of vaporization was applied to calculate the values at other temperatures. However, in many cases the enthalpy of vaporization was not known very well and this might result in a large uncertainty as well.

In the present work, we updated the aerosol module SORGAM with temperature dependent parameterizations of SOA formation considering published chamber experiments of relevance (indicated by the suffix T), we included the SOA formation pathways from isoprene oxidation (indicated by I), and the degradation products initiated by  $\text{NO}_3$  oxidation were also considered (N). The stepwise updated SORGAM-TIN module was first validated in a box model framework with chamber experiments and then tested in the EURAD-IM platform in a case study. Model and measurement comparison of

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the SOA concentrations at a rural location near Rotterdam (the Netherlands) further proved the success of such updates under real European atmospheric conditions.

## 2 Methodology

### 2.1 The current aerosol module in EURAD-IM

5 Within EURAD-IM, the aerosol dynamics such as nucleation, condensation, coagulation, diffusion, sedimentation, aerosol-cloud interaction, etc. are simulated by the Modal Aerosol Dynamics Model for Europe (MADE) (Ackermann et al., 1998). However, the initial MADE aerosol chemistry module only treated inorganic ions and water. To consider the formation of secondary organic aerosols, a new module, SORGAM  
10 was developed and implemented into MADE (Schell et al., 2001). In SORGAM, both anthropogenic and biogenic hydrocarbons are first oxidized by oxidants like OH, NO<sub>3</sub> and O<sub>3</sub>. The mass transfer from gas to particle phase is then driven by the gas/particle partitioning of the low volatility oxidation products formed in gas phase. According to the available chamber experimental results, the fractional aerosol coefficient is used for  
15 the oxidations of alkanes and anthropogenic terminal and internal alkenes, while the two-product model is applied for other reactive organic gases including aromatics and terpenes.

For the two-product model approach (Odum et al., 1996) used in SORGAM, the produced SOA from oxidation of certain hydrocarbon precursor is expressed by

$$20 \quad Y = \frac{M_0 \alpha_1 K_1}{1 + K_1 M_0} + \frac{M_0 \alpha_2 K_2}{1 + K_2 M_0} \quad (1)$$

where  $Y$  is the mass fractional yield of SOA,  $M_0$  is the total aerosol organic mass concentration,  $\alpha_i$  is the mass stoichiometric coefficient, and  $K_i$  is a partitioning coefficient. The subscripts 1 and 2 in Eq. (1) denote the two categories of oxidation products, normally one is assumed to be semi-volatile and the other is attributed to low-

volatile condensable vapors. In the original SORGAM, the mass stoichiometric coefficients  $\alpha_i$  are set to be constants, and the influence of temperature is solely introduced through partitioning coefficients  $K_i$  whose temperature dependence is expressed using the Clausius–Clapeyron equation (Pankow, 1987).

In this work, we updated the SORGAM model to SORGAM-TIN in the following five aspects (a schematic view see Fig. 1): (1) enthalpies of vaporization were updated according to recent study (Saathoff et al., 2009); (2) the influence of aerosol water content on the SOA yield were considered by a simplified RH dependence correction; (3) unified temperature dependent functions of mass stoichiometric coefficients and partitioning coefficients for the two-product approach were derived according to available chamber experiments; (4) the SOA produced from isoprene oxidation products was considered; (5) the SOA produced from the degradation products of biogenic VOCs by  $\text{NO}_3$  was considered as well.

## 2.2 The update of enthalpies of vaporization

The SOA two product species in MADE/SORGAM and their molecular weights and enthalpies of vaporization were listed in Table 1. The enthalpies of vaporization of every product from all precursors were the same in the original SORGAM (Schell et al., 2001). The enthalpies of vaporization of the products of  $\alpha$ -pinene and limonene were updated according to Saathoff et al. (2009); molecular weights and enthalpies of vaporization of the products of isoprene were taken from Carlton et al. (2010) and Kleindienst et al. (2009).

## 2.3 The humidity correction

For the parameterization of SOA formation, the effect of humidity was considered as well in this study. According to Prisle et al. (2010), the activity suppression in the Raoult's law region was a linear function of relative humidity (RH) for any organic

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compound  $i$ :

$$\tau_i(\text{RH}) = 1 - \zeta_i \cdot \text{RH} \quad (2)$$

$$C_i^*(\text{RH}) = \tau_i C_i^*(\text{dry}) \quad (3)$$

- 5 where  $\tau_i$ ,  $\zeta_i$ ,  $C_i^*$  and RH denote organic activity coefficient, hydrophilicity coefficient, saturation mass concentration and relative humidity, respectively. Based on  $\alpha$ -pinene ozonolysis, Prisle et al. (2010) estimated the range of  $\zeta$  varying from 0.33 to 0.67, and within the error limits nearly all the results could be represented by  $\zeta = 0.5$ . Therefore in this study,  $\zeta$  was set to 0.5 for all organic compounds included in SORGAM-TIN.
- 10 Since  $C_i^*$  is the reciprocal of the partitioning coefficient  $K_i$ ,  $K_i(\text{RH})$  can be expressed as  $K_i(\text{dry})/(1 - \zeta \cdot \text{RH})$ . Equation (1) can thus be rearranged to:

$$Y = \frac{M_0 \alpha_1 K_1 / (1 - \zeta \cdot \text{RH})}{1 + K_1 M_0 / (1 - \zeta \cdot \text{RH})} + \frac{M_0 \alpha_2 K_2 / (1 - \zeta \cdot \text{RH})}{1 + K_2 M_0 / (1 - \zeta \cdot \text{RH})} \quad (4)$$

## 2.4 The temperature dependent functions of mass stoichiometric coefficients and partitioning coefficients

15 It was found that the original parameterization of  $\alpha_i$  and  $K_i$  in SORGAM could not describe the available chamber experimental results appropriately, which were published in recent years. With newly available laboratory data collected and reviewed, we deduced new parameterizations of  $\alpha_i$  and  $K_i$  with sophisticated non-linear optimization software packages such as Lsqnonlin, Eureqa (Schmidt and Lipson, 2009) and NI2sol (Dennis et al., 1981). In Fig. 2, the methodological framework of the parameterization was depicted, which included three major steps.

20 Firstly, the selected chamber experimental data were grouped into different temperature bins to deduce the relation between parameters and the temperature in each bin. According to equation Eq. (1), experimentally determined  $Y$  and  $M_0$  are used as input data while the other four parameters ( $\alpha_1$ ,  $K_1$ ,  $\alpha_2$ ,  $K_2$ ) were deduced

in each temperature bin with Lsqnonlin (an optimization tool in Matlab to solve non-linear least squares problems). The results of this step could be summarized as:  $Y_{T_i} = F(M_0, \alpha_1(T_i), \alpha_2(T_i), K_1(T_i), K_2(T_i))$ .

In the second step, we further established the relationship between temperature and the four parameters. With Eureka, a software tool for detecting equations and hidden mathematical relationships in data (Schmidt and Lipson, 2009), we derived intermediate functions relating the stoichiometric coefficients  $\alpha_i$  or partitioning coefficients  $K_i$  with temperature:

$$\alpha_1 = F_1(T); \quad \alpha_2 = F_2(T); \quad K_1 = F_3(T); \quad K_2 = F_4(T).$$

The functions  $F_i$  derived in the second step were tested in a box model framework. However, the model-measurement inter comparison results were not as satisfying as expected (see Sect. 3.2), due to the different algorithm utilized in Eureka and SORGAM. Therefore, in a third step, the localized parameterization of the functions  $F_i$  for SORGAM has to be further deduced. This kind of optimization was achieved with the NI2sol model, which is an adaptive nonlinear least-squares algorithm. Herein, the residual vector, i.e. the differences of simulated and observed values, is a continuously differentiable function  $R(x) = (R_1(x), \dots, R_n(x))^T$  of  $p$  parameters  $x = (x_1, \dots, x_p)^T$ . Vector  $x$  was introduced in the temperature functions like  $F_i(T, x)$ , and the initial value of  $x$  was set as 1.0. NI2sol attempts to find a parameter vector  $x^*$  that minimizes the sum of squares function  $F(x) = \sum R_i(x)^2$  (Dennis et al., 1981). In this way, the final functions  $F_i(T, x^*)$  with best compatibility for SORGAM were obtained.

Due to the limited amount of chamber experiments for toluene, the temperature dependence functions for the parameterizations of toluene were extrapolated from those of m-xylene. In a study by Takekawa et al. (2003) both toluene and m-xylene were investigated under similar experimental conditions. We subdivided these experiments into two temperature groups, namely, 283 K and 303 K. By assuming that the oxidation products of m-xylene and toluene had the same partitioning coefficient  $K_i$ , we thus derived mass stoichiometric coefficient of toluene  $\alpha_{\text{toluene}} = x \cdot \alpha_{\text{m-xylene}}$ ,  $x > 1$  (Vivanco

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et al., 2011). We then solved  $x$  by Eq. (1) for both temperature groups according to available chamber experiments (Takekawa et al., 2003). Linear interpolation between the two temperatures was performed to get the function for factor  $x$ :

$$x = (T - 283) \cdot 0.12 + 1.2 \quad (5)$$

## 2.5 The mechanistic updates

In the original SORGAM, SOA produced from isoprene was not considered. In contrast, as outlined in recent laboratory and field studies (e.g. Claeys et al., 2004; Kroll et al., 2006; Henze and Seinfeld, 2006; Paulot et al., 2009), isoprene could be a more important precursor of SOA than previously thought. Nevertheless, the laboratory results about the SOA yield from isoprene were still limited (Chhabra et al., 2010; Dommen et al., 2006; Kroll et al., 2006), so there are not enough data available to derive temperature dependent functions of  $\alpha$  and  $K$ . Instead, a constant  $\alpha$  valid at 295 K determined by the OH oxidation of isoprene were adapted according to Henze and Seinfeld (2006).



This set of parameters shown in Eq. (6) are widely used in CTMs, including GOES-Chem and CMAQ (Henze and Seinfeld, 2006; Carlton et al., 2010). In addition, the reaction of isoprene and ozone was not included, since this reaction channel would not produce SOA (Pye et al., 2010).

In the original SORGAM, the SOA yields from reactions of  $\text{NO}_3$  with terpenes are assumed to be zero according to Griffin et al. (1999) and the reaction of  $\text{NO}_3$  with isoprene is not considered. In this study, oxidation of  $\alpha$ -pinene, limonene and isoprene by  $\text{NO}_3$  radical was taken into account according to the current state of knowledge (Pye et al., 2010). The partitioning coefficients  $K_i$  for  $\alpha$ -pinene and limonene were calculated with the temperature functions in Table 3. The mass stoichiometric coefficients  $\alpha_i$  for the reaction of isoprene with  $\text{NO}_3$  were set to the largest value experimentally



determined by Ng et al. (2008):



The mass stoichiometric coefficients  $\alpha_i$  for  $\alpha$ -pinene and limonene were referred to the  $\text{NO}_3$  experiments as shown by Table 4 of Griffin et al. (1999). In those experiments, an extremely high aerosol yield ( $\alpha = 1$ ) was determined for the  $\text{NO}_3$  oxidation of bicyclic terpenes. To be compatible with the two-product model applied in SORGAM, it was assumed there were two products with the same stoichiometric coefficients of 0.5:



The mechanistic update of SORGAM module with respect to the  $\text{NO}_3$  oxidation in this study is in line with the current development of GEOS-Chem (Pye et al., 2010; Yu, 2011) and CMAQ (Carlton et al., 2010).

## 3 Results and discussions

### 3.1 Unified temperature parameterized functions – $\alpha(T)$ and $K(T)$

According to recent chamber studies, the SOA yields were found to vary significantly for different parent hydrocarbons and experiment conditions such as temperature. Since data are relative abundant to characterize the temperature dependence for  $\alpha$ -pinene, limonene and m-xylene, chamber studies with similar experimental conditions were used to derive unified temperature parameterized functions for  $\alpha$  and  $K$ . In this study, we grouped the available ozonolysis experiments for  $\alpha$ -pinene, and limonene, and photo-oxidation (OH initiated) experiments for m-xylene (Chen et al., 2010; Cocker et al., 2001a,b; Griffin et al., 1999; Hoffmann et al., 1997; Kamens et al., 1999; Odum et al., 1996; Offenberg et al., 2007; Pathak et al., 2007a,b; Presto et al., 2005, 2006;

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Saathoff et al., 2009; Sarwar et al., 2007; Takekawa et al., 2003; Warren et al., 2009; Yu et al., 1999) as described in Table 2. Due to the fact that records of  $\alpha$ -pinene chamber data were more abundant than others, we were able to define 14 temperature bins ranging from 243 to 320 K. Limonene data, available in the temperature range of 253 to 313 K, was divided into 6 bins. Only 4 bins were used for m-xylene, since the temperature range of experimental data was quite limited (283 to 303 K).

Following the methodology described in Sect. 2.4, we derived the unified temperature parameterized functions  $\alpha(T)$  for mass stoichiometric coefficient ( $\alpha$ ) and  $K(T)$  for gas-particle partitioning coefficient ( $K$ ) of  $\alpha$ -pinene, limonene and m-xylene. We limited the form of the functions as the fundamental arithmetic in Eureka to make them more tractable. On the other hand, the accuracy must be considered as well. If a linear function could not reflect the relation between temperature and a parameter, the order of the function increased, but not higher than 3. In SORGAM-TIN, the derived  $\alpha(T)$  and  $K(T)$  are extended to all oxidation pathways of  $\alpha$ -pinene, limonene and m-xylene assuming that similar oxidation products were produced from OH, NO<sub>3</sub> and O<sub>3</sub>.

To evaluate the effect of the newly derived  $\alpha(T)$  and  $K(T)$ , the temperature dependence of the ratio of the updated SOA yield against the original yield ( $Y_{\text{updated}}/Y_{\text{original}}$ ) for  $\alpha$ -pinene, limonene and m-xylene are further analyzed and displayed in Fig. 3. Limited by the available experimental data, the derived  $\alpha(T)$  and  $K(T)$  is only suitable from 283 K to 304 K. In case the temperature exceeds the scales, the  $\alpha(T)$  and  $K(T)$  will be set as  $\alpha(283)$ ,  $K(283)$  and  $\alpha(304)$ ,  $K(304)$  respectively. In Fig. 3, the yields are calculated based on Eq. 1, and three  $M_0$  values (5, 10, 50  $\mu\text{g m}^{-3}$ ) are selected for different cases. For  $\alpha$ -pinene, updated SOA yields are higher at low  $M_0$  concentration ( $Y_{\text{updated}}/Y_{\text{original}}$  up to 1.80 and 1.37 with  $M_0 = 5$  and 10  $\mu\text{g m}^{-3}$  at lower temperature, respectively). However, when  $M_0$  is high, the updated yields are lower at higher temperatures ( $> 286$  K). For limonene, systematic higher SOA yields are shown in the updated model, with the ratio of  $Y_{\text{updated}}/Y_{\text{original}}$  up to 2.87, 2.24 and 1.62 at lower temperature, based on different  $M_0$  load, respectively. For m-xylene, the updated SOA yields are higher ( $Y_{\text{updated}}/Y_{\text{original}}$  up to 2.43, 3.00 and 3.52 at different  $M_0$ ) at lower temperatures

(< 295 K) while lower ( $Y_{\text{updated}}/Y_{\text{original}}$  down to about 0.24) at higher temperatures (> 295 K).

### 3.2 Performance of the derived $\alpha(T)$ and $K(T)$ in chamber experiments

We further validated the derived  $\alpha(T)$  and  $K(T)$  by comparing the modeled results against available chamber experiments. Original SORGAM model and SORGAM-TIN were employed to simulate the SOA formed by  $\alpha$ -pinene, limonene and m-xylene, and the SOA by isoprene was also tested in the SORGAM-TIN.

In Fig. 4, the original model usually overestimates SOA concentration for all three precursors; but for these precursors, the simulated SOA concentrations are closest to the observation data by SORGAM-TIN. Due to fewer data points available from only one experiment at a fixed temperature, the simulation of isoprene SOA performs very well.

To evaluate the performance of each model quantitatively, normalized mean bias (NMB) and normalized mean errors (NME) were calculated by the following formulae:

$$\text{NMB} = \frac{\sum_{i=1}^n (C_i^s - C_i^o)}{\sum_{i=1}^n C_i^o} \times 100\%, \quad \text{NME} = \frac{\sum_{i=1}^n |C_i^s - C_i^o|}{\sum_{i=1}^n C_i^o} \times 100\%,$$

where  $C_i^s$  and  $C_i^o$  are simulated and observed SOA concentration respectively.

The statistical results of NMB, NME and correlation coefficient (CORREL) calculated for the two models are listed in Table 4. For  $\alpha$ -pinene, the NME decrease from 74.1 % for the original SORGAM to 27.2 % for SORGAM-TIN. For limonene, the simulation result of the original model is good, with NME of 20.6 %, while it is better to use optimized functions in the SORGAM-TIN, with NME of 17.3 %. The optimized functions dramatically reduce the NME for m-xylene (17.4 %), compared to the result of the original model. The result for isoprene SOA with the optimized parameters indicates its good performance. For all precursors, the correlation coefficients between observed and simulated values for the optimized functions are higher than those for the original

model ( $R$  as 0.802 vs. 0.770 for  $\alpha$ -pinene; 0.842 vs. 0.802 for limonene; 0.933 vs. 0.640 for m-xylene). According to the results in Table 4, it was clear that the simulated SOA concentration by the SORGAM-TIN agreed well with the observation in chamber experiments.

We took  $\alpha$ -pinene as an example to show the influence of temperature. In Fig. 5, the temperature test cases are computed with  $\pm 5$  K. Although the temperature influences both reaction rate and partitioning coefficient, it can be seen that the higher SOA concentrations are simulated at low temperature and the lower at high temperature, with the change of 9 % and -6 % in average, respectively. It is noticed that when the mass of SOA is low, the influence of temperature is higher, because with a low  $M_0$ , a small change of SOA concentration may lead to a large rate of change.

### 3.3 The application of SORGAM-TIN in EURAD-IM

#### 3.3.1 The campaign in Cabauw

Within the frame of the EUCAARI project, a comprehensive field campaign – IMPACT (Intensive Measurement Period At Cabauw Tower) campaign was conducted in May 2008 in Cabauw, the Netherlands. The measurement site, namely, the Cabauw Experimental Site for Atmospheric Research (CESAR), is located in the center of the Netherlands (51.972° N, 4.926° E).

During the campaign, the temperature in Cabauw was around 283 to 298 K, and the average temperature was higher in the first half part than that in the second one. The prevailing wind direction was easterly wind of continental origin induced by an anticyclonic regime, while it changed to northerly wind during 15–17 May, due to the passage of a frontal system. Strong precipitation occurred on 16 and 17 May, and another precipitation process took place at the end of the month.

More details about the IMPACT campaign in particular the synoptic conditions can be found in Hamburger et al. (2011). In this work, the observed SOA concentrations were estimated by the summed concentrations of the semi-volatile OOA (SV-OOA) and

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low-volatile OOA (LV-OOA) factors from AMS measurements (Mensah et al., 2012) followed by a positive matrix factorization (PMF) analysis (Docherty et al., 2008). In previous studies, Docherty et al. (2008) calculated the concentration of SOA by some widely used methods, such as EC and CO tracer methods, water soluble organic carbon content, chemical mass balance of organic molecular markers and PMF of AMS data. The comparison of predicted SOA/OA ratios by these methods showed that the ratio by AMS-PMF method was the median and very close to the average of the results of these methods. Therefore the results derived from AMS-PMF method were considered to be a good observational estimation of SOA concentrations. However, the AMS-PMF method also contains two major uncertainties: one is the accuracy of the AMS measurement itself which is mainly limited by the uncertainty of the applied collection efficiency (Canagaratna et al., 2007); the other is the uncertainty of the subsequent PMF analysis (Ulbrich et al., 2009). With respect to the Caubaw dataset, the applied collection efficiency of AMS measurement was validated against a SMPS instrument; the inorganic and organic compositions measured by AMS were inter compared against MARGA and TD-PTR-MS, respectively (Mensah et al., 2012). In general, reasonable good agreements between AMS and MARGA, between AMS and TD-PTR-MS were achieved.

### 3.3.2 EURAD-IM modeling system

The EURAD-IM used in this study is able to simulate the physical and chemical evolution of trace gases and aerosols in the troposphere and lower stratosphere. The meteorological fields needed by EURAD-IM were computed using MM5 Version 3 (Grell et al., 1994). The RACM-MIM-GK chemical mechanism (Stockwell et al., 1997; Geiger et al., 2003; Karl et al., 2006; Lu et al., 2012) was used to describe the atmospheric gas phase chemistry. The aerosol module is MADE containing SORGAM. The spatial distribution of anthropogenic emissions was derived from a high-resolution emission inventory (Kuenen et al., 2011) for the year 2007 based on the TNO emission inventory for 2005 (Denier van der Gon et al., 2010). Biogenic emissions were taken from

Symeonidis et al. (2008) or calculated using the parameterizations of Lamb et al. (1993) for regions where the data is not available in Symeonidis et al. (2008). A horizontal resolution of 45 km was chosen, the atmosphere was divided into 23 vertical layers, and the height of the near surface layer was about 35 m.

Figure 6 shows the comparisons of observed and simulated ozone, sulfate and nitrate. The simulated results can capture the trends of the three species, with the NME of 29.2 % for ozone, 42.0 % for sulfate and 51.0 % for nitrate. The large errors on 16 May might be caused by the failure to simulate the precipitation intensity at Cabauw. In general, the good performances indicate that EURAD-IM does well in the simulation of the photochemical processes and inorganic species.

### 3.3.3 The simulation of SOA with SORGAM-TIN in Cabauw

Comparisons of simulated SOA concentrations by both the SORGAM and SORGAM-TIN and the observed SOA concentrations as represented by the sum of LV-OOA and SV-OOA from AMS measurements are shown in Fig. 7. In Fig. 7a, the observed SOA concentrations were strongly underestimated by SORGAM, while in general were reasonably reproduced by SORGAM-TIN.

Following the division of the generic synoptic system as discussed in Hamburger et al. (2011); the model-measurement inter comparison are further studied for two selected periods: (1) the first half of May (1–15 May), was controlled by subsiding air masses due to an anticyclone event, while (2) the second half of May (18–31 May), was dominated by a westerly frontal passage. Precipitation occurred on 16 and 17 May.

Benelux states always experienced the most serious air pollution over Europe in both periods as characterized by the airborne measurements and ground measurements at Cabauw (Hamburger et al., 2011). Interestingly, the observed SOA concentrations were much better reproduced by SORGAM-TIN for the first period than that of the second period. The difference between these two periods were that the air was quite stagnant in the first period while constant westerly inflow took place during the second

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period. Additionally, the EURAD-IM model also overestimated the temperature at the second period. Therefore, the underestimation of the observed SOA concentration in the second period may be attributed to the deficits of the transport term or temperature corrections of  $K$  and  $\alpha$ . Of the first period, apart from the general good agreements, the SORGAM-TIN had strongly overestimated the observed SOA concentrations on 9 May. The failure of SORGAM-TIN for 16 May can be attributed to a general failure of the EURAD-IM system to simulate time, location, and intensity of precipitation. This general failure could also be found of the large model-measurement discrepancy of  $O_3$ , sulfate and nitrate (Fig. 6). However, the reason for 9 May is probably due to errors of the SORGAM-TIN framework since the other parameters were well simulated in EURAD-IM system.

To separate the contribution of the newly added SOA parameterizations, an accumulative plot of the SOA concentrations modeled by different updates (Table 5 to summarize the sensitivity studies tested with the EURAD-IM system) are shown in Fig. 7b.

The original SORGAM can only predict 20 % of the observed SOA concentrations. The application of both  $\alpha(T)$  and  $K(T)$  enhanced the simulated SOA concentration up to 27 % of the observation, with an increment of 7 %. Moreover, the addition of isoprene SOA further increased the SOA concentration by 15 % of the observation on average. At last, including  $NO_3$  oxidation as SOA formation pathways leads to an average increment of SOA by 32 % of the observation. Overall, the modeled SOA concentration by SORGAM-TIN on average reproduced 88 % of the observation.

Figure 8 further shows the averaged diurnal variations of the observed and modeled SOA concentrations by SORGAM, SORGAM-TIN and the three important mechanistic updates as summarized in Table 5, including the effect of each improvement. The temperature functions usually increase the SOA concentration in the daytime, which even decrease the concentration during the early morning. Adding isoprene won't change the diurnal variation of the SOA concentration, which increases the concentration on average. Employing  $NO_3$  pathways immensely increases the SOA concentration during



night time since the  $\text{NO}_3$  chemistry often becomes significant at nighttime due to the small NO concentration and missing insolation.

The diurnal variation simulated by the SORGAM-TIN is reasonable, during nighttime. In the afternoon the SORGAM-TIN still strongly underestimates the observations, maybe due to the missing of a parameterization of the successive oxidation processes in the model or the thermodynamic assumption is not applicable under such conditions as discussed by (Perraud et al., 2012). The difference between the real atmosphere and experimental conditions for the data in the laboratory studies used for parameterization surely contributes to deficiencies. Furthermore, as we now recognized, the afternoon OH concentrations might be strongly underestimated by current models at rural or suburban conditions (Hofzumahaus et al., 2009; Lu et al., 2012, 2013). Chamber study with mixed anthropogenic and biogenic precursors showed that the anthropogenic SOA could promote the formation of biogenic SOA (Emanuelsson et al., 2012). The underestimated OH concentrations and the missing interaction of anthropogenic and biogenic SOA could also be at least part of the explanations for such underestimated SOA concentrations.

We compare this study to other similar reports regarding the SOA simulation in Europe. The simulation with the CHIMERE model in Paris, spring 2007, was reported by Sciare et al. (2010), of which it was found the SOA was underestimated possibly due to missing potential sources of SOA in the model such as the aging of POA and oligomerisation. Couvidat et al. (2012) simulated the SOA formation from isoprene oxidation with two kinds of parameterizations (Henze and Seinfeld, 2006; Couvidat and Seigneur, 2011), and found out that the latter (Couvidat and Seigneur, 2011) calculated higher concentrations of isoprene SOA indicating the first one (Henze and Seinfeld, 2006). The former is currently widely used in different CTMs and may underestimate isoprene produced SOA masses then. Bergstroem et al. (2012) used the EMEP model with VBS method to simulate long-term OC over Europe, covering 2002–2007. Different VBS schemes were employed, and the results indicated that considering aging of POA and primary emissions of semi-volatile components could signifi-

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cantly increase the simulated SOA concentrations. By comparing to those other similar model-measurement comparisons, we recognized that the missing of aerosol aging processes, the missing oligomerization process, the missing contributions from primary emitted semi-volatile organic compounds (SVOCs) and the parameterization of isoprene SOA formation process could be possibilities of the afternoon SOA underestimations with SORGAM-TIN as well.

Overall the normalized mean error for the updated model in this period was 32.7 %, which was similar with the results in other studies by PMCAMx (Tsimpidi et al., 2010; Murphy et al., 2010), much better than the NME of 80.1 % by the original SORGAM.

## 4 Conclusions

As a continuous effort of EURAD-IM development, we extensively updated the SORGAM module in EURAD-IM in this study. The new module, namely, SORGAM-TIN, received three major modifications: (1) temperature dependent functions of SOA yields, (2) photo oxidation products from isoprene and (3) NO<sub>3</sub> initiated oxidation products from biogenic VOCs. To derive the temperature dependent functions of SOA yields, we summarized the recent published chamber experimental results for  $\alpha$ -pinene, limonene and m-xylene and set them into a series of temperature interval defined bins. Optimization software Lsqnonlin, Eureqa and NI2sol were employed to deduce temperature dependent functions in three steps through nonlinear optimizations. In addition, the influences of relative humidity, temperature and the anthropogenic VOCs representation about toluene SOA formation were also considered. However, they only show marginal contributions to SOA simulations. The SORGAM-TIN was shown to perform well in reproducing chamber experiments, with normalized mean errors less than 30 % in the box model. In contrast, the normalized mean error for the original SORGAM module is much worse especially for  $\alpha$ -pinene (NME=74 %) and m-xylene (NME=70 %).

Both SORGAM-TIN and SORGAM were further tested in chemical transport model EURAD-IM and model results were compared to ambient observations at a rural site

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near Rotterdam, during IMPACT campaign in May 2008. The observed SOA concentrations were represented by summed OOA concentrations extracted from measurement results of an AMS instrument. The original SORGAM model strongly underestimated the observed concentration of SOA which only explained 20 % of the observed concentrations. In contrast, the simulated SOA concentration with SORGAM-TIN could explain on average 88 % of the observed values. The largely improved model results were mainly contributed to the implementation of temperature dependent  $\alpha(T)$  and  $K(T)$ , the photo oxidation of isoprene and the  $\text{NO}_3$  initiated oxidation of biogenic VOCs, contributing 7 %, 15 % and 32 % of the observations, respectively.

However, there are still significant daytime measurement-model discrepancies for SOGRAM-TIN model results, especially for the afternoon hours. Details of the remaining discrepancies remain speculative and still need to be explored in future, when more laboratory and field data are available. In principle, in SORGAM-TIN, considerations of aerosol aging process (Rudich et al., 2007) are missing, and the SOA formation is more likely to be kinetically determined by irreversible uptake processes instead of thermodynamic equilibrium processes (Perraud et al., 2012). In addition, the production rates of the oxidation products during daytime are of course uncertain as well due to a general underestimation of ambient OH concentrations reported for forest, rural and suburban areas (Stone et al., 2012). Moreover, according to our model tests performed herein, the reactions of the  $\text{NO}_3$  radical with biogenic VOCs are found to be important for SOA formation. More laboratory and field experiments on this aspect are therefore required in the future. Oligomerization processes, mixed precursors, multiphase and multigenerational reactions should be included in future work as well.

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*Acknowledgements.* We acknowledge financial support by the National Natural Science Foundation of China (Major Program: 21190052 & Innovative Research Group: 41121004), Strategic Priority Research Program of the Chinese Academy of Sciences (grant no. XDB05010500) and by EU-project PEGASOS (grant no. 265307). Discussions with Zheng, M. are gratefully acknowledged.

## References

- Ackermann, I. J., Hass, H., Memmesheimer, M., Ebel, A., Binkowski, F. S., and Shankar, U.: Modal aerosol dynamics model for Europe: development and first applications, *Atmos. Environ.*, 32, 2981–2999, 1998.
- Bergström, R., Denier van der Gon, H. A. C., Prévôt, A. S. H., Yttri, K. E., and Simpson, D.: Modelling of organic aerosols over Europe (2002–2007) using a volatility basis set (VBS) framework: application of different assumptions regarding the formation of secondary organic aerosol, *Atmos. Chem. Phys.*, 12, 8499–8527, doi:10.5194/acp-12-8499-2012, 2012.
- Canagaratna, M. R., J. T. Jayne, Canagaratna, M. R., Jayne, J. T., Jimenez, J. L., Allan, J. D., Alfarra, M. R., Zhang, Q., Onasch, T. B., Drewnick, F., Coe, H., Middlebrook, A., Delia, A., Williams, L. R., Trimborn, A. M., Northway, M. J., DeCarlo, P. F., Kolb, C. E., Davidovits, P., and Worsnop, D. R.: Chemical and microphysical characterization of ambient aerosols with the aerodyne aerosol mass spectrometer, *Mass Spectrom. Rev.* 26, 185–222, 2007.
- Cappa, C. D., Zhang, X., Loza, C. L., Craven, J. S., Yee, L. D., and Seinfeld, J. H.: Application of the Statistical Oxidation Model (SOM) to Secondary Organic Aerosol formation from photooxidation of C12 alkanes, *Atmos. Chem. Phys.*, 13, 1591–1606, doi:10.5194/acp-13-1591-2013, 2013.
- Carlton, A. G., Bhawe, P. V., Napelenok, S. L., Edney, E. D., Sarwar, G., Pinder, R. W., Pouliot, G. A., and Houyoux, M.: Model Representation of Secondary Organic Aerosol in CMAQv4.7, *Environ. Sci. Tech.*, 44, 8553–8560, 2010.
- Chen, X. and Hopke, P. K.: A chamber study of secondary organic aerosol formation by limonene ozonolysis, *Indoor Air*, 20, 320–328, 2010.

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- Chervenkov, H. and Jakobs, H.: Dust storm simulation with regional air quality model – problems and results, *Atmos. Environ.*, 45, 3965–3976, 2011.
- Chhabra, P. S., Flagan, R. C., and Seinfeld, J. H.: Elemental analysis of chamber organic aerosol using an aerodyne high-resolution aerosol mass spectrometer, *Atmos. Chem. Phys.*, 10, 4111–4131, doi:10.5194/acp-10-4111-2010, 2010.
- Chung, S. H. and Seinfeld, J. H.: Global distribution and climate forcing of carbonaceous aerosols, *J. Geophys. Res.-Atmos.*, 107, 2002.
- Claeys, M., Graham, B., Vas, G., Wang, W., Vermeylen, R., Pashynska, V., Cafmeyer, J., Guyon, P., Andreae, M. O., Artaxo, P., and Maenhaut, W.: Formation of secondary organic aerosols through photooxidation of isoprene, *Science*, 303, 1173–1176, 2004.
- Cocker, D. R., Clegg, S. L., Flagan, R. C., and Seinfeld, J. H.: The effect of water on gas-particle partitioning of secondary organic aerosol. Part I: alpha-pinene/ozone system, *Atmos. Environ.*, 35, 6049–6072, 2001a.
- Cocker, D. R., Mader, B. T., Kalberer, M., Flagan, R. C., and Seinfeld, J. H.: The effect of water on gas-particle partitioning of secondary organic aerosol: II. m-xylene and 1,3,5-trimethylbenzene photooxidation systems, *Atmos. Environ.*, 35, 6073–6085, 2001b.
- Colette, A., Granier, C., Hodnebrog, Ø., Jakobs, H., Maurizi, A., Nyiri, A., Rao, S., Amann, M., Bessagnet, B., D'Angiola, A., Gauss, M., Heyes, C., Klimont, Z., Meleux, F., Memmesheimer, M., Mieville, A., Rouil, L., Russo, F., Schucht, S., Simpson, D., Stordal, F., Tampieri, F., and Vrac, M.: Future air quality in Europe: a multi-model assessment of projected exposure to ozone, *Atmos. Chem. Phys.*, 12, 10613–10630, doi:10.5194/acp-12-10613-2012, 2012.
- Couvidat, F. and Seigneur, C.: Modeling secondary organic aerosol formation from isoprene oxidation under dry and humid conditions, *Atmos. Chem. Phys.*, 11, 893–909, doi:10.5194/acp-11-893-2011, 2011.
- Couvidat, F., Debry, E., Sartelet, K., and Seigneur, C.: A hydrophilic/hydrophobic organic ((HO)-O-2) aerosol model: development, evaluation and sensitivity analysis, *J. Geophys. Res.-Atmos.*, 117, 2012.
- Denier van der Gon, H., Visschedijk, A., Kuenen, J., van der Brugh, H., Dröge, R., and Schaap, M.: High resolution European emission grids for anthropogenic sources for the years 2003–2007, EGU 2010 special session Monitoring Atmospheric Composition and Climate, *Geophys. Res. Abstr.*, 12, EGU2010-10800-1, 2010.
- Dennis, J. E., Gay, D. M., and Welsch, R. E.: Algorithm 573 – NL2SOL – an adaptive non-linear least-squares algorithm E4, *ACM T. Math. Software*, 7, 369–383, 1981.

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Dommen, J., Metzger, A., Duplissy, J., Kalberer, M., Alfara, M. R., Gascho, A., Weingartner, E., Prevot, A. S. H., Verheggen, B., and Baltensperger, U.: Laboratory observation of oligomers in the aerosol from isoprene/NO<sub>x</sub> photooxidation, *Geophys. Res. Lett.*, 33, L13805 doi:10.1029/2006gl026523, 2006.

Donahue, N. M., Robinson, A. L., Stanier, C. O., and Pandis, S. N.: Coupled partitioning, dilution, and chemical aging of semivolatile organics, *Environ. Sci. Tech.*, 40, 2635–2643, 2006.

Elbern, H., Strunk, A., Schmidt, H., and Talagrand, O.: Emission rate and chemical state estimation by 4-dimensional variational inversion, *Atmos. Chem. Phys.*, 7, 3749–3769, doi:10.5194/acp-7-3749-2007, 2007.

Emanuelsson, E. U., Hallquist, M., Kristensen, K., Glasius, M., Bohn, B., Fuchs, H., Kammer, B., Kiendler-Scharr, A., Nehr, S., Rubach, F., Tillmann, R., Wahner, A., Wu, H.-C., and Mentel, Th. F.: Formation of anthropogenic secondary organic aerosol (SOA) and its influence on biogenic SOA properties, *Atmos. Chem. Phys. Discuss.*, 12, 20311–20350, doi:10.5194/acpd-12-20311-2012, 2012.

Fry, J. L., Kiendler-Scharr, A., Rollins, A. W., Wooldridge, P. J., Brown, S. S., Fuchs, H., Dubé, W., Mensah, A., dal Maso, M., Tillmann, R., Dorn, H.-P., Brauers, T., and Cohen, R. C.: Organic nitrate and secondary organic aerosol yield from NO<sub>3</sub> oxidation of  $\beta$ -pinene evaluated using a gas-phase kinetics/aerosol partitioning model, *Atmos. Chem. Phys.*, 9, 1431–1449, doi:10.5194/acp-9-1431-2009, 2009.

Fry, J. L., Kiendler-Scharr, A., Rollins, A. W., Brauers, T., Brown, S. S., Dorn, H.-P., Dubé, W. P., Fuchs, H., Mensah, A., Rohrer, F., Tillmann, R., Wahner, A., Wooldridge, P. J., and Cohen, R. C.: SOA from limonene: role of NO<sub>3</sub> in its generation and degradation, *Atmos. Chem. Phys.*, 11, 3879–3894, doi:10.5194/acp-11-3879-2011, 2011.

Fu, T.-M., Jacob, D. J., Wittrock, F., Burrows, J. P., Vrekoussis, M., and Henze, D. K.: Global budgets of atmospheric glyoxal and methylglyoxal, and implications for formation of secondary organic aerosols, *J. Geophys. Res.-Atmos.*, 113, D15303 doi:10.1029/2007jd009505, 2008.

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- Geiger, H., Barnes, I., Bejan, J., Benter, T., and Spittler, M.: The tropospheric degradation of isoprene: an updated module for the regional atmospheric chemistry mechanism, *Atmos. Environ.*, 37, 1503–1519, 2003.
- 5 Grell, G. A., Dudhia, J., and Stauffer, D. R.: A description of the fifth generation Penn State/NCAR mesoscale model (MM5); NCAR TN-398+IA, technical report, National Center for Atmospheric Research, Boulder, Colorado, 1994.
- Griffin, R. J., Cocker, D. R., Flagan, R. C., and Seinfeld, J. H.: Organic aerosol formation from the oxidation of biogenic hydrocarbons, *J. Geophys. Res.-Atmos.*, 104, 3555–3567, 1999.
- 10 Guenther, A., Karl, T., Harley, P., Wiedinmyer, C., Palmer, P. I., and Geron, C.: Estimates of global terrestrial isoprene emissions using MEGAN (Model of Emissions of Gases and Aerosols from Nature), *Atmos. Chem. Phys.*, 6, 3181–3210, doi:10.5194/acp-6-3181-2006, 2006.
- Guenther, A., Zimmerman, P., Klinger, L., Greenberg, J., Ennis, C., Davis, K., Pollock, W., Westberg, H., Allwine, G., and Geron, C.: Estimates of regional natural volatile organic compound fluxes from enclosure and ambient measurements, *J. Geophys. Res.-Atmos.*, 101, 1345–1359, 1996.
- 15 Hallquist, M., Wenger, J. C., Baltensperger, U., Rudich, Y., Simpson, D., Claeys, M., Dommen, J., Donahue, N. M., George, C., Goldstein, A. H., Hamilton, J. F., Herrmann, H., Hoffmann, T., Iinuma, Y., Jang, M., Jenkin, M. E., Jimenez, J. L., Kiendler-Scharr, A., Maenhaut, W., McFiggans, G., Mentel, Th. F., Monod, A., Prévôt, A. S. H., Seinfeld, J. H., Surratt, J. D., Szmigielski, R., and Wildt, J.: The formation, properties and impact of secondary organic aerosol: current and emerging issues, *Atmos. Chem. Phys.*, 9, 5155–5236, doi:10.5194/acp-9-5155-2009, 2009.
- 20 Hamburger, T., McMeeking, G., Minikin, A., Birmili, W., Dall'Osto, M., O'Dowd, C., Flentje, H., Henzing, B., Junninen, H., Kristensson, A., de Leeuw, G., Stohl, A., Burkhardt, J. F., Coe, H., Krejci, R., and Petzold, A.: Overview of the synoptic and pollution situation over Europe during the EUCAARI-LONGREX field campaign, *Atmos. Chem. Phys.*, 11, 1065–1082, doi:10.5194/acp-11-1065-2011, 2011.
- 25 Heald, C. L., Jacob, D. J., Park, R. J., Russell, L. M., Huebert, B. J., Seinfeld, J. H., Liao, H., and Weber, R. J.: A large organic aerosol source in the free troposphere missing from current models, *Geophys. Res. Lett.*, 32, L18809, doi:10.1029/2005gl023831, 2005.
- 30 Henze, D. K. and Seinfeld, J. H.: Global secondary organic aerosol from isoprene oxidation, *Geophys. Res. Lett.*, 33, 4, 2006.



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Hoffmann, T., Odum, J. R., Bowman, F., Collins, D., Klockow, D., Flagan, R. C., and Seinfeld, J. H.: Formation of organic aerosols from the oxidation of biogenic hydrocarbons, *J. Atmos. Chem.*, 26, 189–222, 1997.

Hofzumahaus, A., Rohrer, F., Lu, K. D., Bohn, B., Brauers, T., Chang, C. C., Fuchs, H., Holland, F., Kita, K., Kondo, Y., Li, X., Lou, S. R., Shao, M., Zeng, L. M., Wahner, A., and Zhang, Y. H.: Amplified trace gas removal in the troposphere, *Science*, 324, 1702–1704, 2009.

Jimenez, J. L., Canagaratna, M. R., Donahue, N. M., Prevot, A. S. H., Zhang, Q., Kroll, J. H., DeCarlo, P. F., Allan, J. D., Coe, H., Ng, N. L., Aiken, A. C., Docherty, K. S., Ulbrich, I. M., Grieshop, A. P., Robinson, A. L., Duplissy, J., Smith, J. D., Wilson, K. R., Lanz, V. A., Hueglin, C., Sun, Y. L., Tian, J., Laaksonen, A., Raatikainen, T., Rautiainen, J., Vaattovaara, P., Ehn, M., Kulmala, M., Tomlinson, J. M., Collins, D. R., Cubison, M. J., Dunlea, E. J., Huffman, J. A., Onasch, T. B., Alfarra, M. R., Williams, P. I., Bower, K., Kondo, Y., Schneider, J., Drewnick, F., Borrmann, S., Weimer, S., Demerjian, K., Salcedo, D., Cottrell, L., Griffin, R., Takami, A., Miyoshi, T., Hatakeyama, S., Shimono, A., Sun, J. Y., Zhang, Y. M., Dzepina, K., Kimmel, J. R., Sueper, D., Jayne, J. T., Herndon, S. C., Trimborn, A. M., Williams, L. R., Wood, E. C., Middlebrook, A. M., Kolb, C. E., Baltensperger, U., and Worsnop, D. R.: Evolution of organic aerosols in the atmosphere, *Science*, 326, 1525–1529, 2009.

Jonsson, Å. M., Hallquist, M., and Ljungström, E.: The effect of temperature and water on secondary organic aerosol formation from ozonolysis of limonene,  $\Delta^3$ -carene and  $\alpha$ -pinene, *Atmos. Chem. Phys.*, 8, 6541–6549, doi:10.5194/acp-8-6541-2008, 2008.

Kamens, R., Jang, M., Chien, C. J., and Leach, K.: Aerosol formation from the reaction of alpha-pinene and ozone using a gas-phase kinetics aerosol partitioning model, *Environ. Sci. Tech.*, 33, 1430–1438, 1999.

Kamens, R. M., Zhang, H. F., Chen, E. H., Zhou, Y., Parikh, H. M., Wilson, R. L., Galloway, K. E., and Rosen, E. P.: Secondary organic aerosol formation from toluene in an atmospheric hydrocarbon mixture: water and particle seed effects, *Atmos. Environ.*, 45, 2324–2334, 2011.

Karl, M., Dorn, H. P., Holland, F., Koppmann, R., Poppe, D., Rupp, L., Schaub, A., and Wahner, A.: Product study of the reaction of OH radicals with isoprene in the atmosphere simulation chamber SAPHIR, *J. Atmos. Chem.*, 55, 167–187, 2006.

Kavouras, I. G., Mihalopoulos, N., and Stephanou, E. G.: Formation of atmospheric particles from organic acids produced by forests, *Nature*, 395, 683–686, 1998.

- Kesselmeier, J. and Staudt, M.: Biogenic volatile organic compounds (VOC): an overview on emission, physiology and ecology, *J. Atmos. Chem.*, 33, 23–88, 1999.
- Kleindienst, T. E., Lewandowski, M., Offenberg, J. H., Jaoui, M., and Edney, E. O.: The formation of secondary organic aerosol from the isoprene + OH reaction in the absence of NO<sub>x</sub>, *Atmos. Chem. Phys.*, 9, 6541–6558, doi:10.5194/acp-9-6541-2009, 2009.
- Kroll, J. H., Ng, N. L., Murphy, S. M., Flagan, R. C., and Seinfeld, J. H.: Secondary organic aerosol formation from isoprene photooxidation, *Environ. Sci. Tech.*, 40, 1869–1877, 2006.
- Kuenen, J., Denier van der Gon, H., Visschedijk, A., and van der Brugh, H.: High resolution European emission inventory for the years 2003–2007, TNO report TNO-060-UT-2011-00588, Utrecht, 2011.
- Lamb, B., Gay, D., Westberg, H., and Pierce, T.: A biogenic hydrocarbon emission inventory for the USA using a simple forest canopy model, *Atmos. Environ. A Gen.*, 27, 1673–1690, 1993.
- Lane, T. E., Donahue, N. M., and Pandis, S. N.: Simulating secondary organic aerosol formation using the volatility basis-set approach in a chemical transport model, *Atmos. Environ.*, 42, 7439–7451, 2008.
- Lu, K. D., Rohrer, F., Holland, F., Fuchs, H., Bohn, B., Brauers, T., Chang, C. C., Häsel, R., Hu, M., Kita, K., Kondo, Y., Li, X., Lou, S. R., Nehr, S., Shao, M., Zeng, L. M., Wahner, A., Zhang, Y. H., and Hofzumahaus, A.: Observation and modelling of OH and HO<sub>2</sub> concentrations in the Pearl River Delta 2006: a missing OH source in a VOC rich atmosphere, *Atmos. Chem. Phys.*, 12, 1541–1569, doi:10.5194/acp-12-1541-2012, 2012.
- Lu, K. D., Hofzumahaus, A., Holland, F., Bohn, B., Brauers, T., Fuchs, H., Hu, M., Häsel, R., Kita, K., Kondo, Y., Li, X., Lou, S. R., Oebel, A., Shao, M., Zeng, L. M., Wahner, A., Zhu, T., Zhang, Y. H., and Rohrer, F.: Missing OH source in a suburban environment near Beijing: observed and modelled OH and HO<sub>2</sub> concentrations in summer 2006, *Atmos. Chem. Phys.*, 13, 1057–1080, doi:10.5194/acp-13-1057-2013, 2013.
- Mensah, A. A., Holzinger, R., Otjes, R., Tramborn, A., Mentel, Th. F., ten Brink, H., Henzing, B., and Kiendler-Scharr, A.: Aerosol chemical composition at Cabauw, The Netherlands as observed in two intensive periods in May 2008 and March 2009, *Atmos. Chem. Phys.*, 12, 4723–4742, doi:10.5194/acp-12-4723-2012, 2012.
- Monteiro, A., Strunk, A., Carvalho, A., Tchepel, O., Miranda, A. I., Borrego, C., Saavedra, S., Rodriguez, A., Souto, J., Casares, J., Friese, E., and Elbern, H.: Investigating a high ozone episode in a rural mountain site, *Environ. Pollut.*, 162, 176–189, 2012.

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- Murphy, B. N. and Pandis, S. N.: Exploring summertime organic aerosol formation in the eastern United States using a regional-scale budget approach and ambient measurements, *J. Geophys. Res.-Atmos.*, 115, D24216, doi:10.1029/2010JD014418, 2010.
- Naumann, K. H.: COSIMA – a computer program simulating the dynamics of fractal aerosols, *J. Aerosol Sci.*, 34, 1371–1397, 2003.
- Ng, N. L., Kwan, A. J., Surratt, J. D., Chan, A. W. H., Chhabra, P. S., Sorooshian, A., Pye, H. O. T., Crounse, J. D., Wennberg, P. O., Flagan, R. C., and Seinfeld, J. H.: Secondary organic aerosol (SOA) formation from reaction of isoprene with nitrate radicals ( $\text{NO}_3$ ), *Atmos. Chem. Phys.*, 8, 4117–4140, doi:10.5194/acp-8-4117-2008, 2008.
- Odum, J. R., Hoffmann, T., Bowman, F., Collins, D., Flagan, R. C., and Seinfeld, J. H.: Gas/particle partitioning and secondary organic aerosol yields, *Environ. Sci. Tech.*, 30, 2580–2585, 1996.
- Odum, J. R., Jungkamp, T. P. W., Griffin, R. J., Flagan, R. C., and Seinfeld, J. H.: The atmospheric aerosol-forming potential of whole gasoline vapor, *Science*, 276, 96–99, 1997.
- Offenberg, J. H., Lewis, C. W., Lewandowski, M., Jaoui, M., Kleindienst, T. E., and Edney, E. O.: Contributions of toluene and alpha-pinene to SOA formed in an irradiated toluene/alpha-pinene/ $\text{NO}_x$ /air mixture: comparison of results using C-14 content and SOA organic tracer methods, *Environ. Sci. Tech.*, 41, 3972–3976, 2007.
- Pandis, S. N., Paulson, S. E., Seinfeld, J. H., and Flagan, R. C.: Aerosol formation in the photooxidation of isoprene and beta-pinene, *Atmos. Environ. A Gen.*, 25, 997–1008, 1991.
- Pankow, J. F.: Review and comparative-analysis of the theories on partitioning between the gas and aerosol particulate phases in the atmosphere, *Atmos. Environ.*, 21, 2275–2283, 1987.
- Pathak, R. K., Presto, A. A., Lane, T. E., Stanier, C. O., Donahue, N. M., and Pandis, S. N.: Ozonolysis of  $\alpha$ -pinene: parameterization of secondary organic aerosol mass fraction, *Atmos. Chem. Phys.*, 7, 3811–3821, doi:10.5194/acp-7-3811-2007, 2007a.
- Pathak, R. K., Stanier, C. O., Donahue, N. M., and Pandis, S. N.: Ozonolysis of alpha-pinene at atmospherically relevant concentrations: temperature dependence of aerosol mass fractions (yields), *J. Geophys. Res.-Atmos.*, 112, 8, D03201 doi:10.1029/2006jd007436, 2007b.
- Paulot, F., Crounse, J. D., Kjaergaard, H. G., Kurten, A., St Clair, J. M., Seinfeld, J. H., and Wennberg, P. O.: Unexpected epoxide formation in the gas-phase photooxidation of isoprene, *Science*, 325, 730–733, 2009.
- Perraud, V., Bruns, E. A., Ezell, M. J., Johnson, S. N., Yu, Y., Alexander, M. L., Zelenyuk, A., Imre, D., Chang, W. L., Dabdub, D., Pankow, J. F., and Finlayson-Pitts, B. J.: Nonequilibrium

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atmospheric secondary organic aerosol formation and growth, P. Natl. Acad. Sci. USA, 109, 2836–2841, 2012.

Presto, A. A. and Donahue, N. M.: Investigation of alpha-pinene plus ozone secondary organic aerosol formation at low total aerosol mass, Environ. Sci. Tech., 40, 3536–3543, 2006.

5 Presto, A. A., Hartz, K. E. H., and Donahue, N. M.: Secondary organic aerosol production from terpene ozonolysis. 2. Effect of NO<sub>x</sub> concentration, Environ. Sci. Tech., 39, 7046–7054, 2005.

10 Prisle, N. L., Engelhart, G. J., Bilde, M., and Donahue, N. M.: Humidity influence on gas-particle phase partitioning of alpha-pinene + O<sub>3</sub> sary organic aerosol, Geophys. Res. Lett., 37, 5, L01802, doi:10.1029/2009gl041402, 2010.

Pye, H. O. T., Chan, A. W. H., Barkley, M. P., and Seinfeld, J. H.: Global modeling of organic aerosol: the importance of reactive nitrogen (NO<sub>x</sub> and NO<sub>3</sub>), Atmos. Chem. Phys., 10, 11261–11276, doi:10.5194/acp-10-11261-2010, 2010.

15 Qi, L., Nakao, S., Tang, P., and Cocker III, D. R.: Temperature effect on physical and chemical properties of secondary organic aerosol from m-xylene photooxidation, Atmos. Chem. Phys., 10, 3847–3854, doi:10.5194/acp-10-3847-2010, 2010.

20 Rollins, A. W., Kiendler-Scharr, A., Fry, J. L., Brauers, T., Brown, S. S., Dorn, H.-P., Dubé, W. P., Fuchs, H., Mensah, A., Mentel, T. F., Rohrer, F., Tillmann, R., Wegener, R., Wooldridge, P. J., and Cohen, R. C.: Isoprene oxidation by nitrate radical: alkyl nitrate and secondary organic aerosol yields, Atmos. Chem. Phys., 9, 6685–6703, doi:10.5194/acp-9-6685-2009, 2009.

Rudich, Y., Donahue, N. M., and Mentel, T. F.: Aging of organic aerosol: bridging the gap between laboratory and field studies, Ann. Rev. Phys. Chem., 58, 321–352, 2007.

25 Saathoff, H., Naumann, K.-H., Möhler, O., Jonsson, Å. M., Hallquist, M., Kiendler-Scharr, A., Mentel, Th. F., Tillmann, R., and Schurath, U.: Temperature dependence of yields of secondary organic aerosols from the ozonolysis of  $\alpha$ -pinene and limonene, Atmos. Chem. Phys., 9, 1551–1577, doi:10.5194/acp-9-1551-2009, 2009.

Sarwar, G., and Corsi, R.: The effects of ozone/limonene reactions on indoor secondary organic aerosols, Atmos. Environ., 41, 959–973, 2007.

30 Schell, B., Ackermann, I. J., Hass, H., Binkowski, F. S., and Ebel, A.: Modeling the formation of secondary organic aerosol within a comprehensive air quality model system, J. Geophys. Res.-Atmos., 106, 28275–28293, 2001.

Schmidt, M., and Lipson, H.: Distilling free-form natural laws from experimental data, Science, 324, 81–85, 2009.

ACPD

13, 5961–6005, 2013

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- Sciare, J., d'Argouges, O., Zhang, Q. J., Sarda-Estève, R., Gaimoz, C., Gros, V., Beekmann, M., and Sanchez, O.: Comparison between simulated and observed chemical composition of fine aerosols in Paris (France) during springtime: contribution of regional versus continental emissions, *Atmos. Chem. Phys.*, 10, 11987–12004, doi:10.5194/acp-10-11987-2010, 2010.
- 5 Seinfeld, J. H. and Pankow, J. F.: Organic atmospheric particulate material, *Ann. Rev. Phys. Chem.*, 54, 121–140, 2003.
- Seinfeld, J. H., Erdakos, G. B., Asher, W. E., and Pankow, J. F.: Modeling the formation of secondary organic aerosol (SOA). 2. The predicted effects of relative humidity on aerosol formation in the alpha-pinene-, beta-pinene-, sabinene-, Delta(3)-Carene-, and cyclohexene-ozone systems, *Environ. Sci. Tech.*, 35, 1806–1817, 2001.
- 10 Stanier, C. O., Donahue, N. M., and Pandis, S. N.: Parameterization of secondary organic aerosol mass fractions from smog chamber data, *Atmos. Environ.*, 42, 2276–2299, 2008.
- Stockwell, W. R., Kirchner, F., Kuhn, M., and Seinfeld, S.: A new mechanism for regional atmospheric chemistry modeling, *J. Geophys. Res.-Atmos.*, 102, 25847–25879, 1997.
- 15 Stone, D., Whalley, L. K., and Heard, D. E.: Tropospheric OH and HO<sub>2</sub> radicals: field measurements and model comparisons, *Chem. Soc. Rev.*, 41, 6348–6404, 2012.
- Symeonidis, P., Gkantou, A., Melas, D., Yay, O. D., Pouspourika, E., Balis, D., and Poupkou, A.: Development of a computational system for estimating biogenic NMVOCs emissions based on GIS technology, *Atmos. Environ.*, 42, 1777–1789, 2008.
- 20 Takekawa, H., Minoura, H., and Yamazaki, S.: Temperature dependence of secondary organic aerosol formation by photo-oxidation of hydrocarbons, *Atmos. Environ.*, 37, 3413–3424, 2003.
- Tillmann, R., Hallquist, M., Jonsson, Å. M., Kiendler-Scharr, A., Saathoff, H., Iinuma, Y., and Mentel, Th. F.: Influence of relative humidity and temperature on the production of pinonaldehyde and OH radicals from the ozonolysis of  $\alpha$ -pinene, *Atmos. Chem. Phys.*, 10, 7057–7072, doi:10.5194/acp-10-7057-2010, 2010.
- 25 Tsimpidi, A. P., Karydis, V. A., Zavala, M., Lei, W., Molina, L., Ulbrich, I. M., Jimenez, J. L., and Pandis, S. N.: Evaluation of the volatility basis-set approach for the simulation of organic aerosol formation in the Mexico City metropolitan area, *Atmos. Chem. Phys.*, 10, 525–546, doi:10.5194/acp-10-525-2010, 2010.
- 30 Ulbrich, I. M., Canagaratna, M. R., Zhang, Q., Worsnop, D. R., and Jimenez, J. L.: Interpretation of organic components from Positive Matrix Factorization of aerosol mass spectrometric data, *Atmos. Chem. Phys.*, 9, 2891–2918, doi:10.5194/acp-9-2891-2009, 2009.

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Vivanco, M. G., Santiago, M., Martinez-Tarifa, A., Borrás, E., Rodenas, M., Garcia-Diego, C., and Sanchez, M.: SOA formation in a photoreactor from a mixture of organic gases and HONO for different experimental conditions, *Atmos. Environ.*, 45, 708–715, 2011.

Warren, B., Austin, R. L., and Cocker, D. R.: Temperature dependence of secondary organic aerosol, *Atmos. Environ.*, 43, 3548–3555, 2009.

Yu, F.: A secondary organic aerosol formation model considering successive oxidation aging and kinetic condensation of organic compounds: global scale implications, *Atmos. Chem. Phys.*, 11, 1083–1099, doi:10.5194/acp-11-1083-2011, 2011.

Yu, J. Z., Cocker, D. R., Griffin, R. J., Flagan, R. C., and Seinfeld, J. H.: Gas-phase ozone oxidation of monoterpenes: Gaseous and particulate products, *J. Atmos. Chem.*, 34, 207–258, 1999.

Zhang, Q., Jimenez, J. L., Canagaratna, M. R., Allan, J. D., Coe, H., Ulbrich, I., Alfarra, M. R., Takami, A., Middlebrook, A. M., Sun, Y. L., Dzepina, K., Dunlea, E., Docherty, K., DeCarlo, P. F., Salcedo, D., Onasch, T., Jayne, J. T., Miyoshi, T., Shimojo, A., Hatakeyama, S., Takegawa, N., Kondo, Y., Schneider, J., Drewnick, F., Borrmann, S., Weimer, S., Demerjian, K., Williams, P., Bower, K., Bahreini, R., Cottrell, L., Griffin, R. J., Rautiainen, J., Sun, J. Y., Zhang, Y. M., and Worsnop, D. R.: Ubiquity and dominance of oxygenated species in organic aerosols in anthropogenically-influenced Northern Hemisphere midlatitudes, *Geophys. Res. Lett.*, 34, L13801 doi:10.1029/2007GL029979, 2007.

Zhang, X. and Seinfeld, J. H.: A Functional Group Oxidation Model (FGOM) for SOA formation and aging, *Atmos. Chem. Phys. Discuss.*, 12, 32565–32611, doi:10.5194/acpd-12-32565-2012, 2012.

Zhou, Y., Zhang, H. F., Parikh, H. M., Chen, E. H., Rattanavaraha, W., Rosen, E. P., Wang, W. X., and Kamens, R. M.: Secondary organic aerosol formation from xylenes and mixtures of toluene and xylenes in an atmospheric urban hydrocarbon mixture: water and particle seed effects (II), *Atmos. Environ.*, 45, 3882–3890, 2011.

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**Table 1.** SOA two product species in SORGAM module.

| RACM Precursor Classes | Oxidants                                 | SOA Model Classes | MW (g mol <sup>-1</sup> ) | $\Delta H_{\text{vap}}$ (kJ mol <sup>-1</sup> ) |
|------------------------|--|-------------------|---------------------------|---|
| TOL, XYL, CSL          | OH/NO <sub>3</sub>                       | ARO1              | 150                       | 156   |
|                        |  | ARO2              | 150                       | 156   |
| HC<br>OLT, OLI         | OH<br>O <sub>3</sub> /OH/NO <sub>3</sub> | ALK1              | 140                       | 156   |
|                        |  | OLE1              | 140                       | 156   |
|                        |  | API1              | 184                       | 24 <sup>a</sup>                                 |
| API                    | O <sub>3</sub> /OH/NO <sub>3</sub>       | API2              | 184                       | 59 <sup>a</sup>                                 |
|                        |  | LIM1              | 200                       | 25 <sup>a</sup>                                 |
| LIM                    | O <sub>3</sub> /OH/NO <sub>3</sub>       | LIM2              | 200                       | 55 <sup>a</sup>                                 |
|                        |  | ISO1              | 96 <sup>b</sup>           | 39 <sup>c</sup>                                 |
| <i>ISO</i>             | <i>O<sub>3</sub>/OH/NO<sub>3</sub></i>   | ISO2              | 96 <sup>b</sup>           | 39 <sup>c</sup>                                 |

<sup>a</sup> Saathoff et al. (2009).<sup>b</sup> Carlton et al. (2010).<sup>c</sup> Kleindienst et al. (2009).

Italic and bold type: updates of original SORGAM.

(TOL: toluene and less reactive aromatics; XYL: xylene and more reactive aromatics; CSL: cresol and other hydroxyl substituted aromatics; HC: alkanes; OLT: terminal alkenes; OLI: internal alkenes; API:  $\alpha$ -pinene and other cyclic terpenes with one double bond; LIM: limonene and other cyclic diene-terpenes; ISO: isoprene.)

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**Table 2.** Temperature grouping of the ozonolysis experiments of  $\alpha$ -pinene and limonene, and photo-oxidation (OH initiated) experiments of m-xylene for deriving unified temperature parameterized functions for  $\alpha$  and  $K$ .

| $\alpha$ -pinene |     |             | Limonene |     |             | m-xylene |     |             |
|------------------|-----|-------------|----------|-----|-------------|----------|-----|-------------|
| No.              | $T$ | Data points | No.      | $T$ | Data points | No.      | $T$ | Data points |
| 1                | 243 | 3           | 8        | 295 | 19          | 1        | 283 | 3           |
| 2                | 252 | 2           | 9        | 300 | 11          | 2        | 293 | 22          |
| 3                | 273 | 4           | 10       | 301 | 12          | 3        | 298 | 10          |
| 4                | 278 | 7           | 11       | 302 | 23          | 4        | 303 | 3           |
| 5                | 284 | 2           | 12       | 303 | 16          | 5        | 303 | 1           |
| 6                | 288 | 11          | 13       | 313 | 22          | 6        | 313 | 1           |
| 7                | 293 | 14          | 14       | 320 | 5           |          |     |             |

The data in each group was taken from different references, listed as following:

$\alpha$ -pinene: Cocker et al. (2001a); Griffin et al. (1999); Hoffmann et al. (1997); Kamens et al. (1999); Offenberger et al. (2007); Pathak et al. (2007a,b); Presto et al. (2005, 2006); Saathoff et al. (2009); Warren et al. (2009); Yu et al. (1999).

limonene: Saathoff et al. (2009); Sarwar et al. (2007); Chen et al. (2010).

m-xylene: Takekawa et al. (2003); Cocker et al. (2001b); Odum et al. (1996).

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**Table 3a.** The unified temperature parameterized functions  $\alpha(T)$  for mass stoichiometric coefficient ( $\alpha$ ).

| species          | $\alpha_1$   | $\alpha_2$   |
|------------------|--|--|
| $\alpha$ -pinene | $\alpha_1 = 0.03315 + \frac{13.377}{T - 179.17}$                                 | $\alpha_2 = \frac{6186.77}{T} + 0.0659 \cdot T - 40.296$                         |
| limonene         | $\alpha_1 = 2.018e^{-3} \cdot T - 0.3114$  | $\alpha_2 = 3.32 - 0.0106 \cdot T$   |
| m-xylene         | $\alpha_1 = 4.912e^{-5} \cdot T^3 + 12.67 \cdot T - 0.0432 \cdot T^2 - 1237.702$ | $\alpha_2 = 8.180e^{-5} \cdot T^3 + 21.18 \cdot T - 0.0721 \cdot T^2 - 2071.922$ |

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**Table 3b.** The unified temperature parameterized functions  $K(T)$  for gas-particle partitioning coefficient ( $K$ ).

| species          | $K_1$  | $K_2$  |
|------------------|--|--|
| $\alpha$ -pinene | $K_1 = \frac{2.419}{3.658e^{-4}T^2 - 0.181T + 22.35}$                        | $K_2 = \frac{4605.54}{121.175T^2 - 59611.81T + 7319862.5}$                     |
| limonene         | $K_1 = \frac{1000.55}{T - 245.94} - 16.7212$                                 | $K_2 = \frac{227.58}{T - 228.84} - 1.0581$                                     |
| m-xylene         | $K_1 = 2.675e^{-3} \cdot T^3 + 691.21 \cdot T - 2.356 \cdot T^2 - 67553.914$ | $K_2 = 5.952e^{-6} \cdot T^3 + 1.55 \cdot T - 5.256e^{-3} \cdot T^2 - 151.349$ |

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**Table 4.** Evaluation of box model simulations with SORGAM and SORGAM-TIN for all species.

| Species          |        | SORGAM | SORGAM-TIN |
|------------------|--------|--------|------------|
| $\alpha$ -pinene | NMB    | 66.5 % | −0.9 %     |
|                  | NME    | 74.1 % | 27.2 %     |
|                  | CORREL | 0.770  | 0.802      |
| limonene         | NMB    | 13.8 % | −2.4 %     |
|                  | NME    | 20.6 % | 17.3 %     |
|                  | CORREL | 0.802  | 0.842      |
| m-xylene         | NMB    | 68.0 % | −0.3 %     |
|                  | NME    | 70.0 % | 17.4 %     |
|                  | CORREL | 0.640  | 0.933      |
| isoprene         | NMB    | –      | 0.2 %      |
|                  | NME    | –      | 9.2 %      |
|                  | CORREL | –      | 0.989      |

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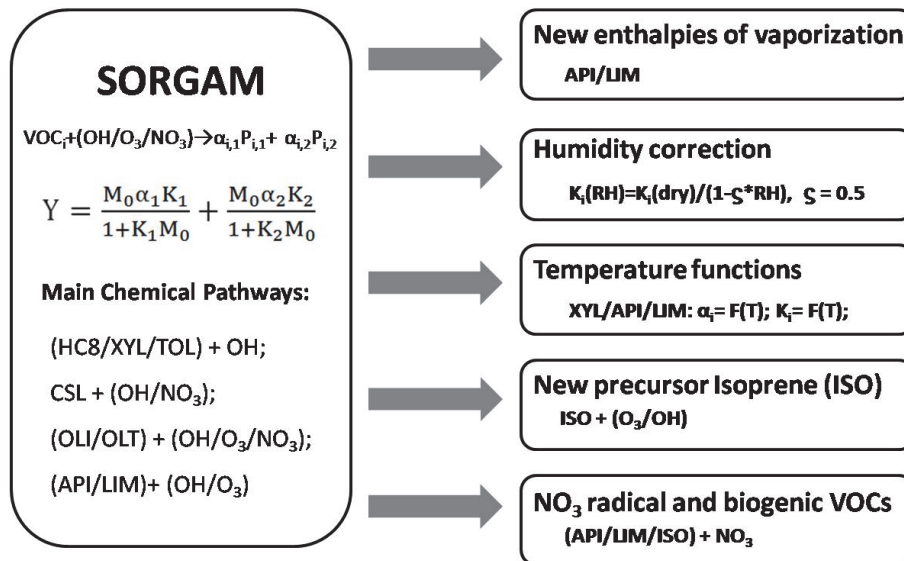
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**Table 5.** The performance of adjusted models.

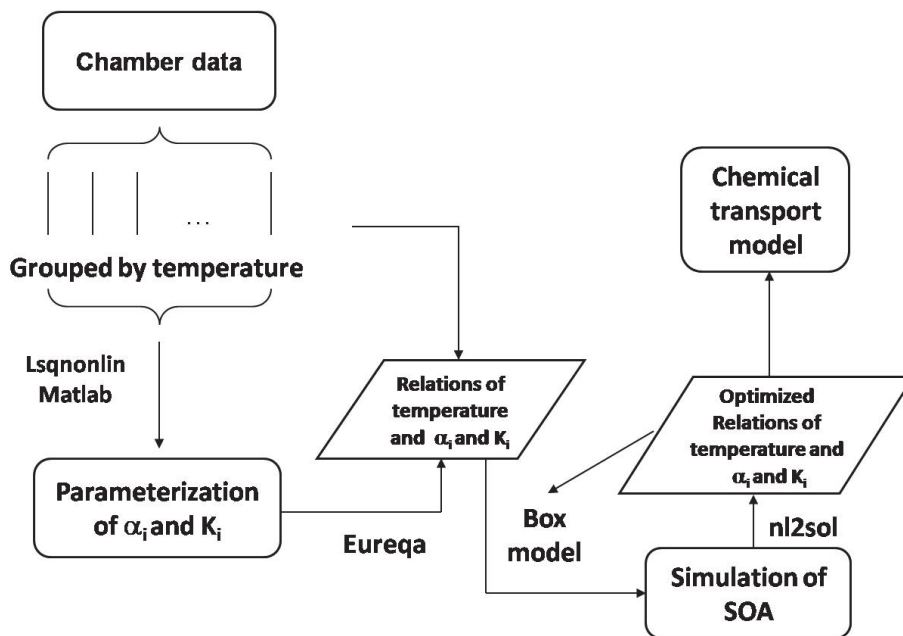
|                 | Simulations/observation | Simulations/original simulation |
|-----------------|-------------------------|---------------------------------|
| Original SORGAM | 0.20                    | 1.00                            |
| SORGAM-T        | 0.27                    | 1.35                            |
| SORGAM-I        | 0.35                    | 1.75                            |
| SORGAM-N        | 0.52                    | 2.60                            |
| SORGAM-TIN      | 0.88                    | 4.40                            |



**Fig. 1.** A schematic view of the updates we applied to the original SORGAM module.

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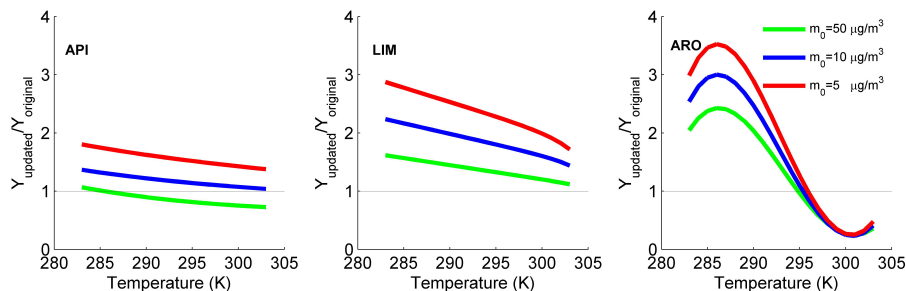
**Fig. 2.** Technical frame of parameterization of mass stoichiometric coefficients and partitioning coefficients. First step: grouping experimental data, second step: fitting parameters in each group, third step: deriving temperature dependent functions, last step: optimizing parameter functions.

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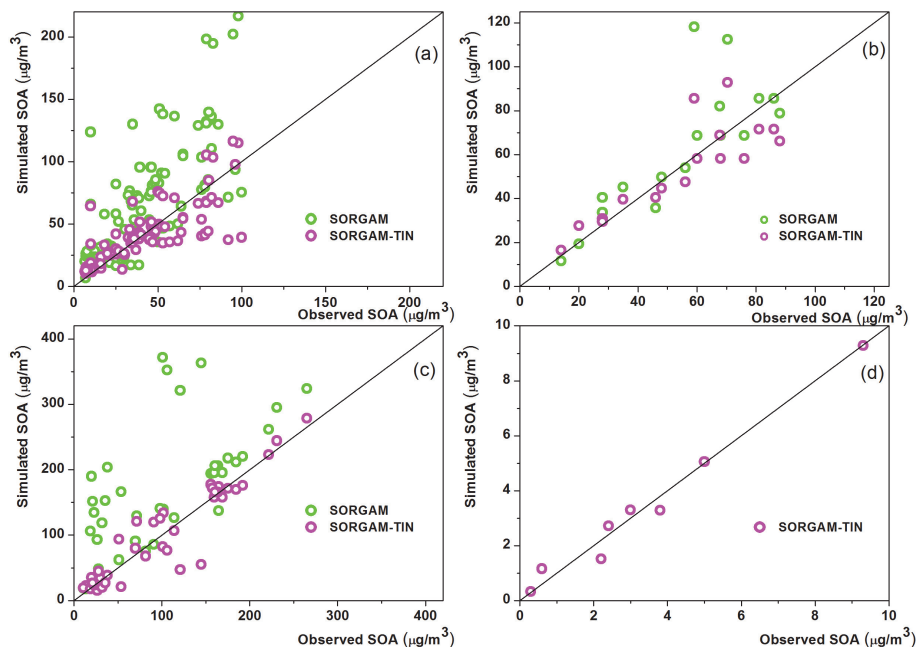


**Fig. 3.** Temperature dependence of the ratio between the updated SOA yields ( $Y_{\text{updated}}$ ) calculated with temperature parameterized  $\alpha$  and  $K$  and the original SOA yields ( $Y_{\text{original}}$ ) used in SORGAM for mechanistic categories API, LIM and ARO for 3 different mass concentrations.

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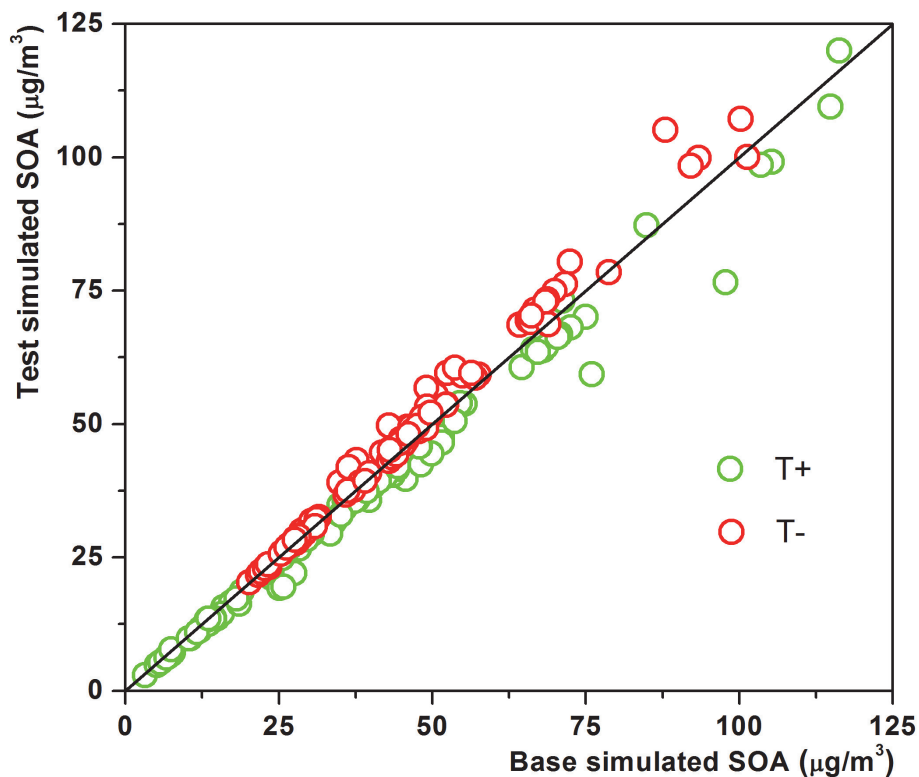

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**Fig. 4.** Comparisons of observed chamber experiment data (x-axis) with original SORGAM (green) and SORGAM-TIN (pink) simulations (y-axis) for **(a):**  $\alpha$ -pinene; **(b):** limonene; **(c):** m-xylene; **(d):** isoprene.

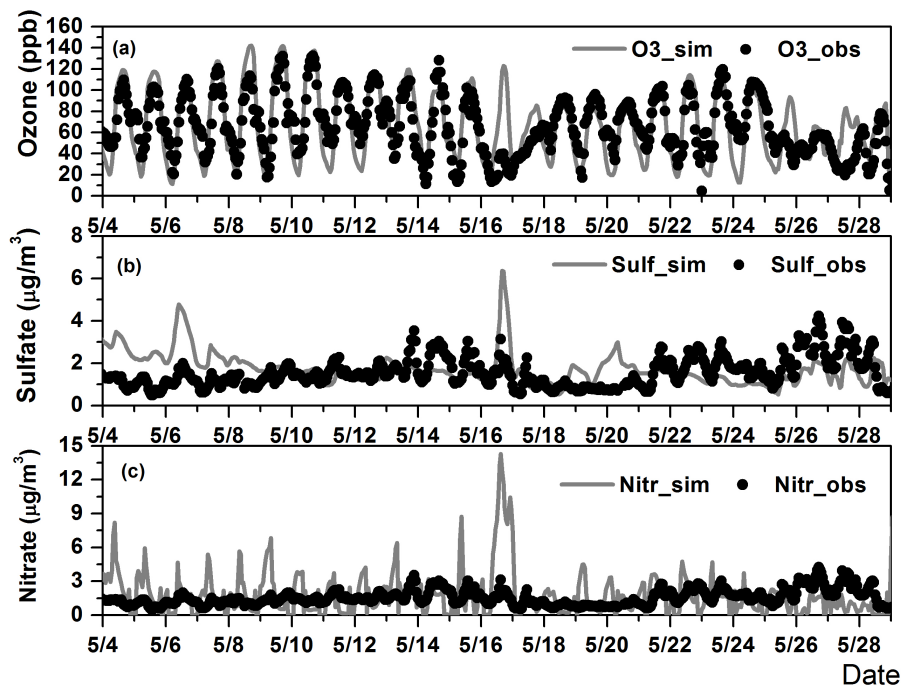
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**Fig. 5.** Box model sensitivity study on temperature for SOA mass produced from  $\alpha$ -pinene. The base case is the same as the final result in Fig. 4, and the test cases are calculated with 5 K higher ( $T+$ ) and 5 K lower ( $T-$ ) temperature.

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**Fig. 6.** Time series of observed (black dots) and simulated (gray line) **(a)** ozone, **(b)** sulfate and **(c)** nitrate in Cabauw May 2008.

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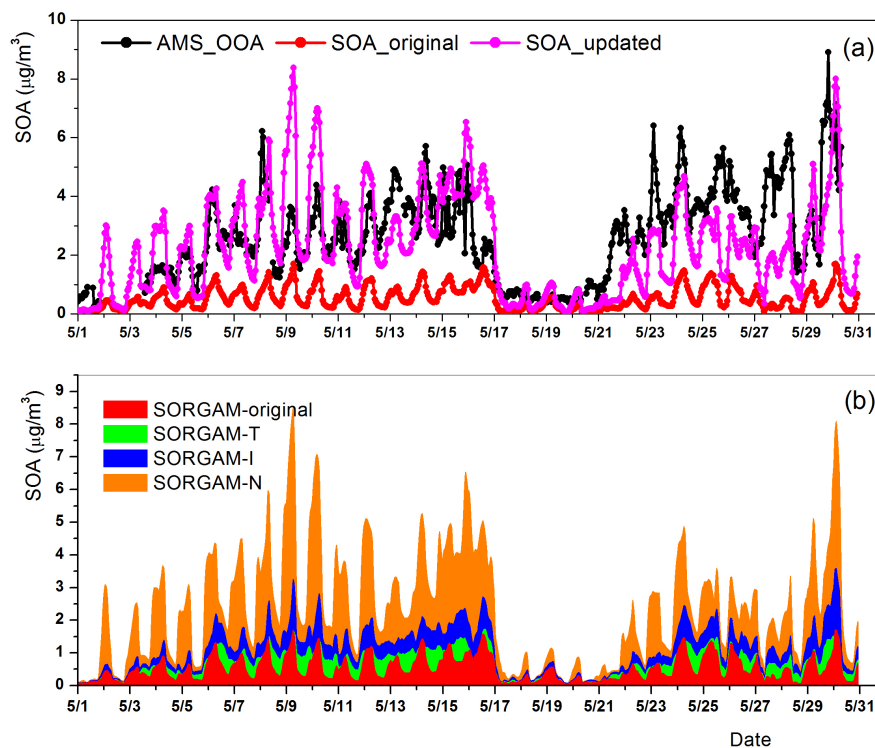
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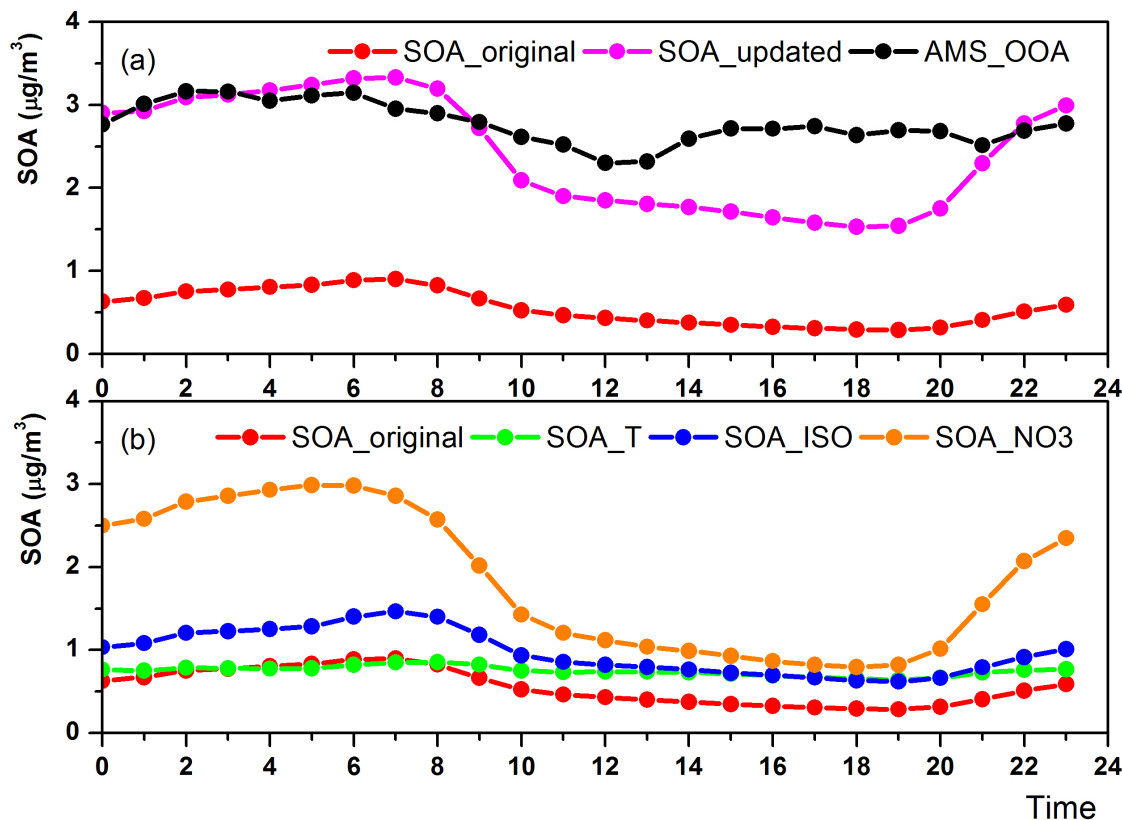


**Fig. 7.** Time series of AMS\_OOA and simulated SOA concentration in Cabauw 2008. **(a)** Black: observation; red: simulation results of original SORGAM; pink: results of SORGAM-TIN. **(b)** Green: increased SOA concentration by the temperature functions; blue: increased SOA concentration by isoprene; orange: increased SOA concentration by  $\text{NO}_3$  chemical pathways.

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**Fig. 8.** Mean diurnal profiles of observed and modeled SOA concentrations in Cabauw from 5–29 May, 2008. **(a)** Black: observation; red: simulation results of original SORGAM; pink: results of the updated SOA module. **(b)** Green: original SORGAM with temperature functions; blue: original SORGAM with isoprene; orange: original SORGAM with NO<sub>3</sub> chemical pathways.